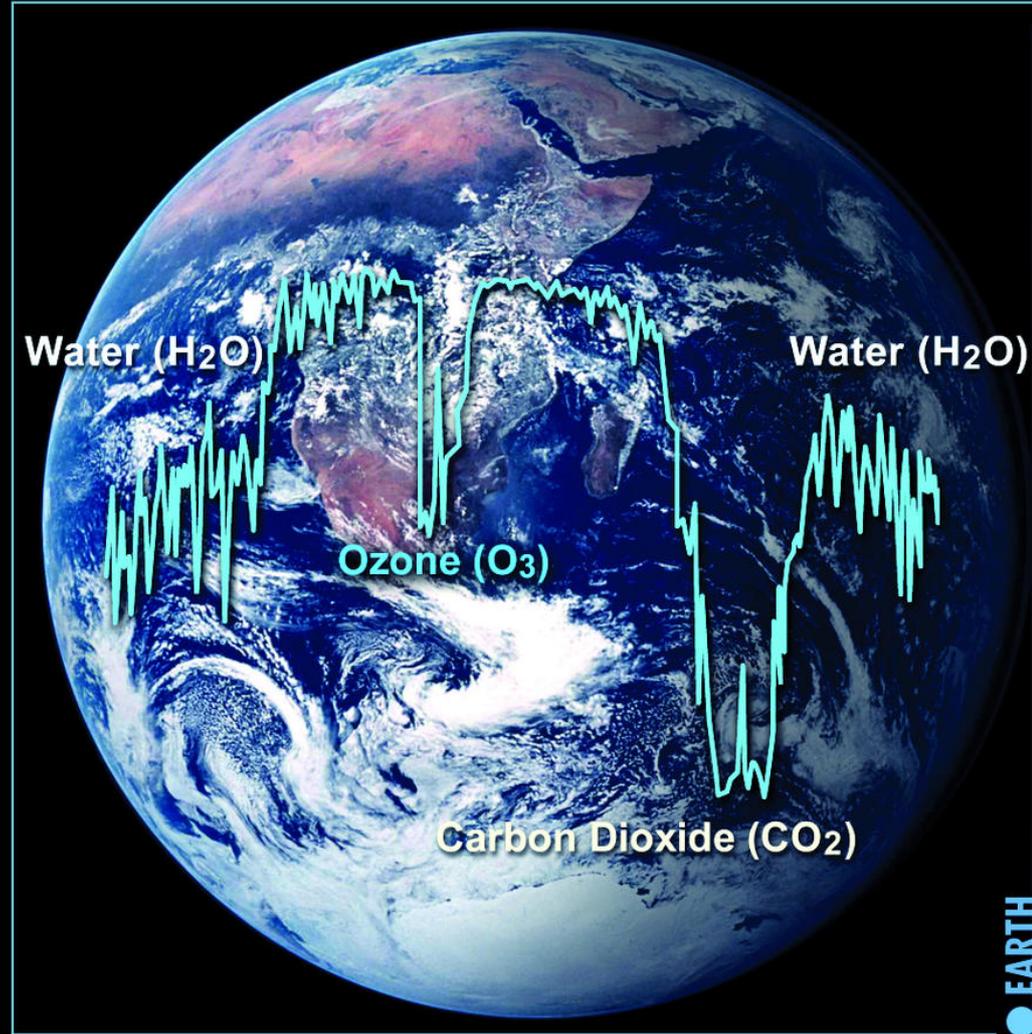
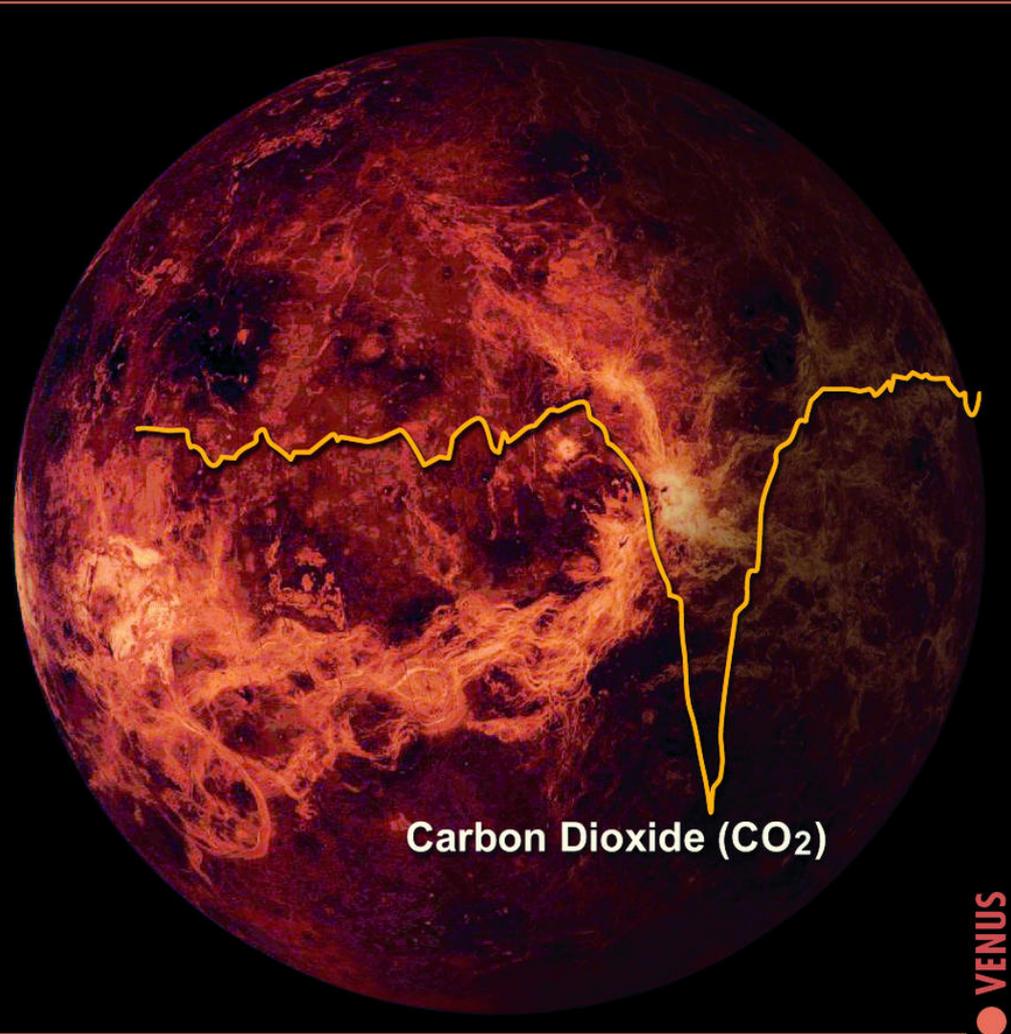


# Planetary Habitability



Stephen Kane

# Topics

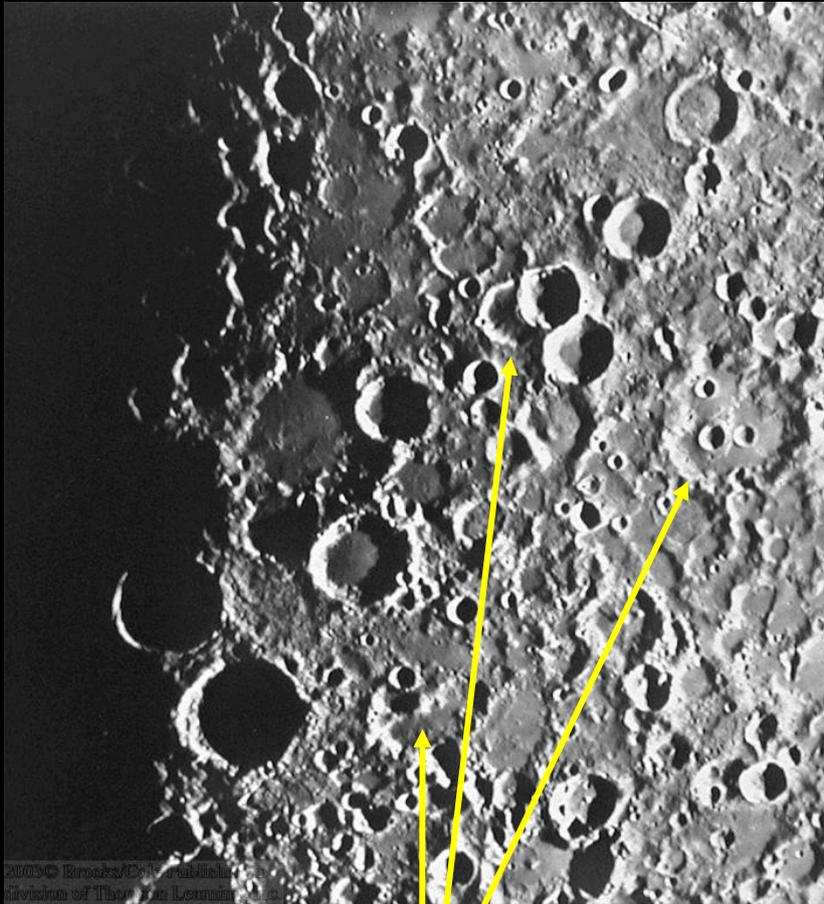
- **Lecture 1 - Introduction**
- **Lecture 2 - Habitability Factors**
- **Lecture 3 - Stars**
- **Lecture 4 - Planetary Atmospheres**
- **Lecture 5 - Planetary Interiors**
- **Lecture 6 - Planetary Energy Balance**
- **Lecture 7 - Habitable Zone I**
- **Lecture 8 - Habitable Zone II**
- **Lecture 9 - Earth as a Living Planet**
- **Lecture 10 - Mars**
- **Lecture 11 - Icy Moons**
- **Lecture 12 - Venus**
- **Lecture 13 - Mercury & the Moon**
- **Lecture 14 - The Role of Giant Planets**
- **Lecture 15 - Stellar Influences**
- **Lecture 16 - Magnetic Fields**
- **Lecture 17 - Milankovitch Cycles**
- **Lecture 18 - Geological Cycles**
- **Lecture 19 - The Next Steps**
- **Lecture 20 - Summary/Discussion**

# **The Moon & Mercury**

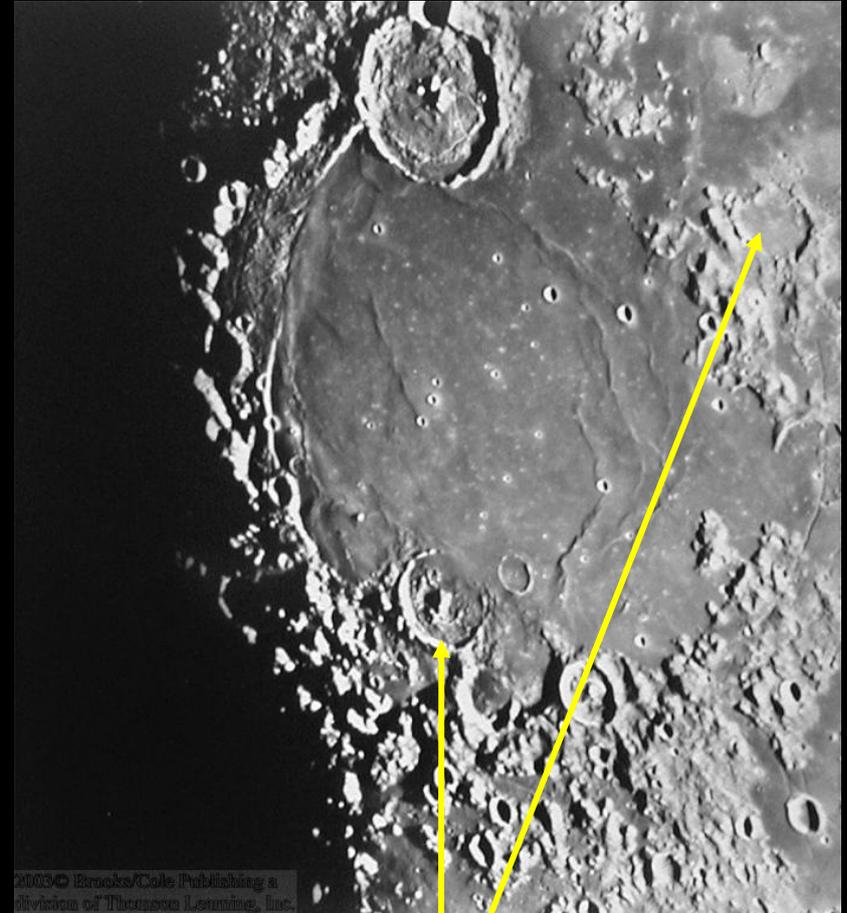
- **Lunar impacts**
- **Formation of the moon**
- **Mercury properties**
- **Surface and interior**
- **Atmosphere and missions**
- **Importance of the Moon & Mercury**

# The Moon: Highlands

Saturated with craters



Older craters partially  
obliterated by more recent  
impacts

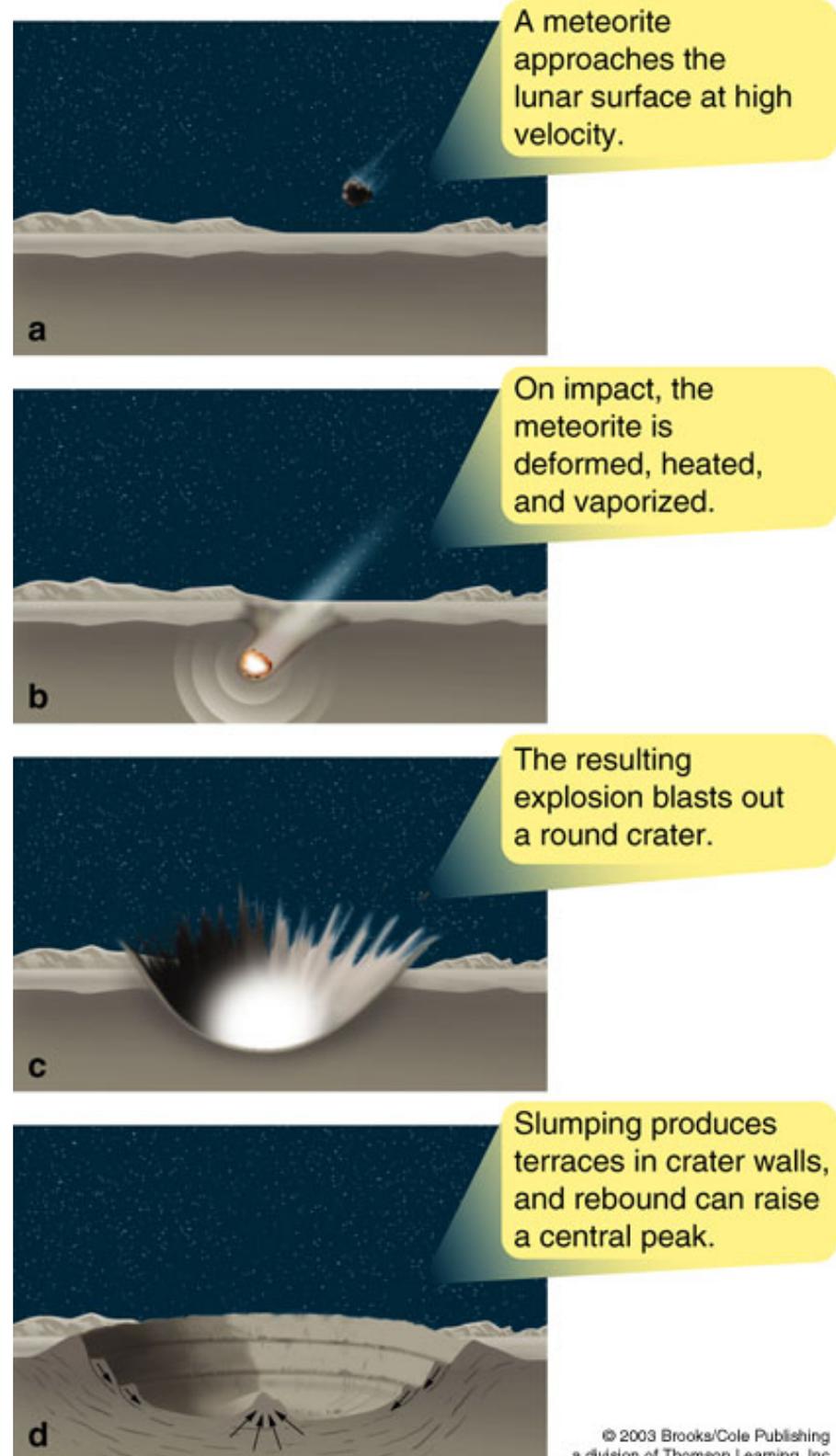
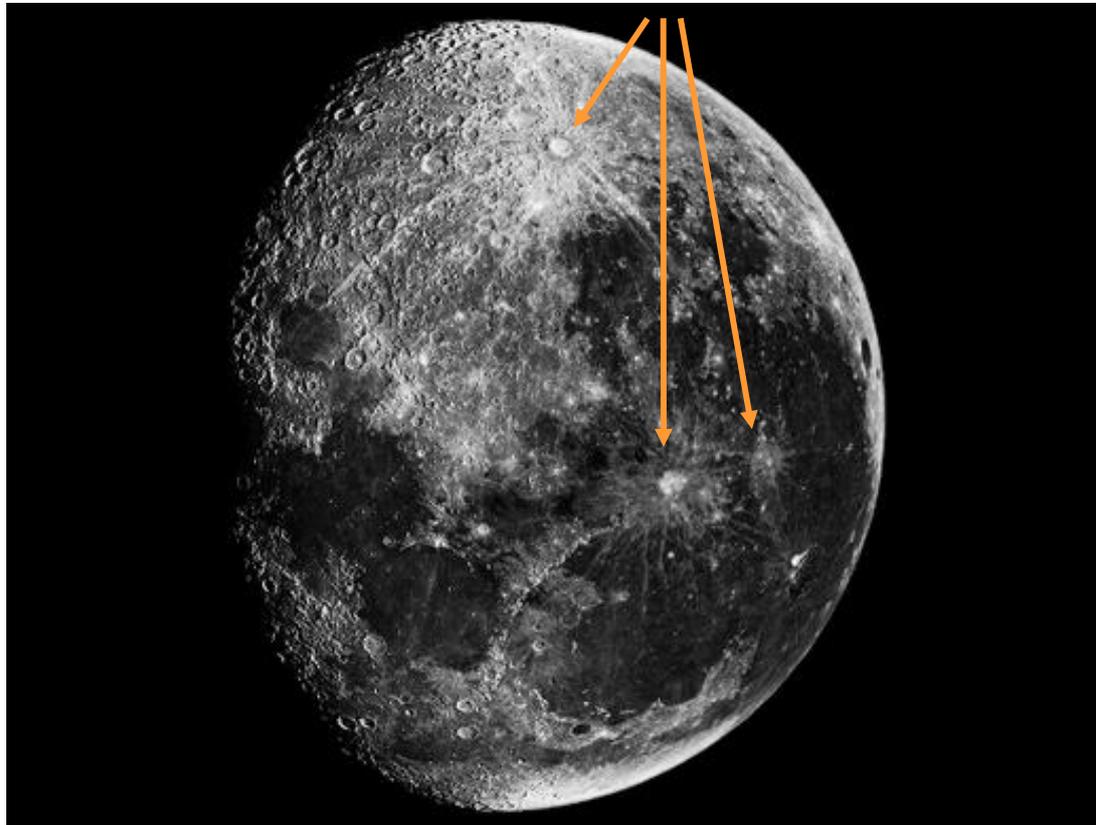


... or flooded by lava  
flows

# Impact Cratering

Impact craters on the moon can be seen easily even with small telescopes.

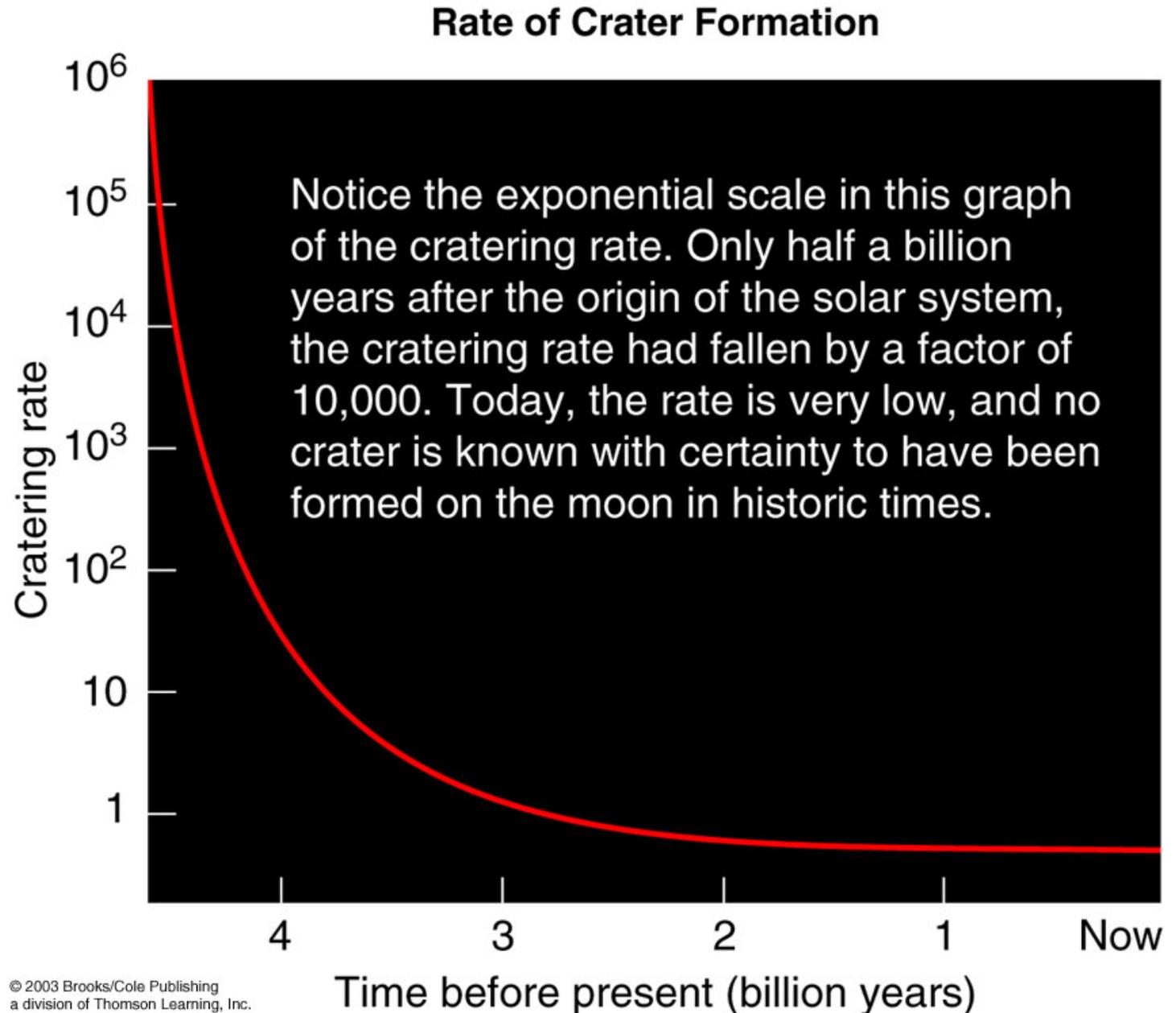
Ejecta from the impact can be seen as bright rays originating from young craters



# History of Impact Cratering

Rate of impacts due to interplanetary bombardment decreased rapidly after the formation of the solar system.

Most craters seen on the Moon's (and Mercury's) surface were formed within the first  $\sim \frac{1}{2}$  billion years.



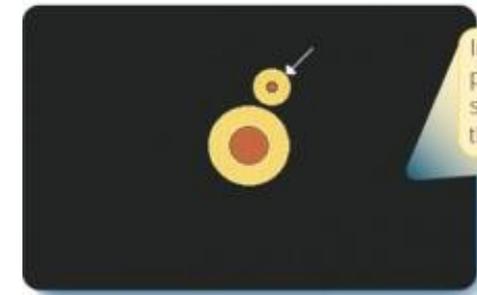


# Modern Theory of Formation of the Moon

## The Large-Impact Hypothesis

- Impact heated material enough to melt it
  - consistent with “sea of magma”
- Collision after differentiation of Earth’s interior
  - Different chemical compositions of Earth and moon
  - (in particular: The moon does not have a large iron core.)

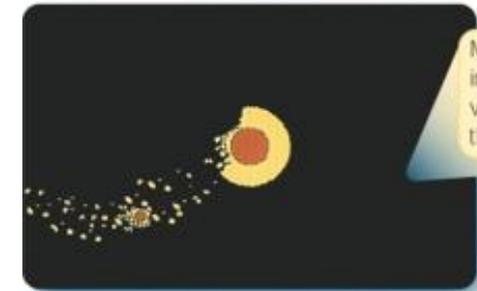
The Large-Impact Hypothesis



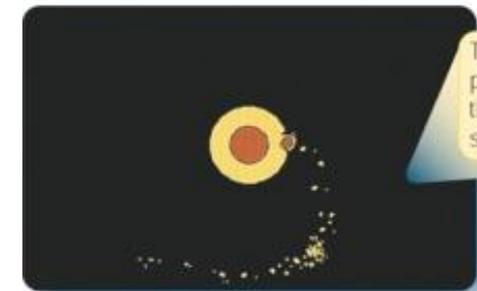
In this model a planetesimal the size of Mars strikes the proto-Earth.



Both bodies have differentiated before the impact.



Material ejected is iron poor, and volatiles are lost in the vacuum of space.



The iron core of the planetesimal falls into the proto-Earth and sinks.



The moon begins to form from volatile-poor rock containing little iron.

# The History of the Moon

Moon is small; low mass → rapidly cooling off; small escape velocity → no atmosphere → unprotected against meteorite impacts.

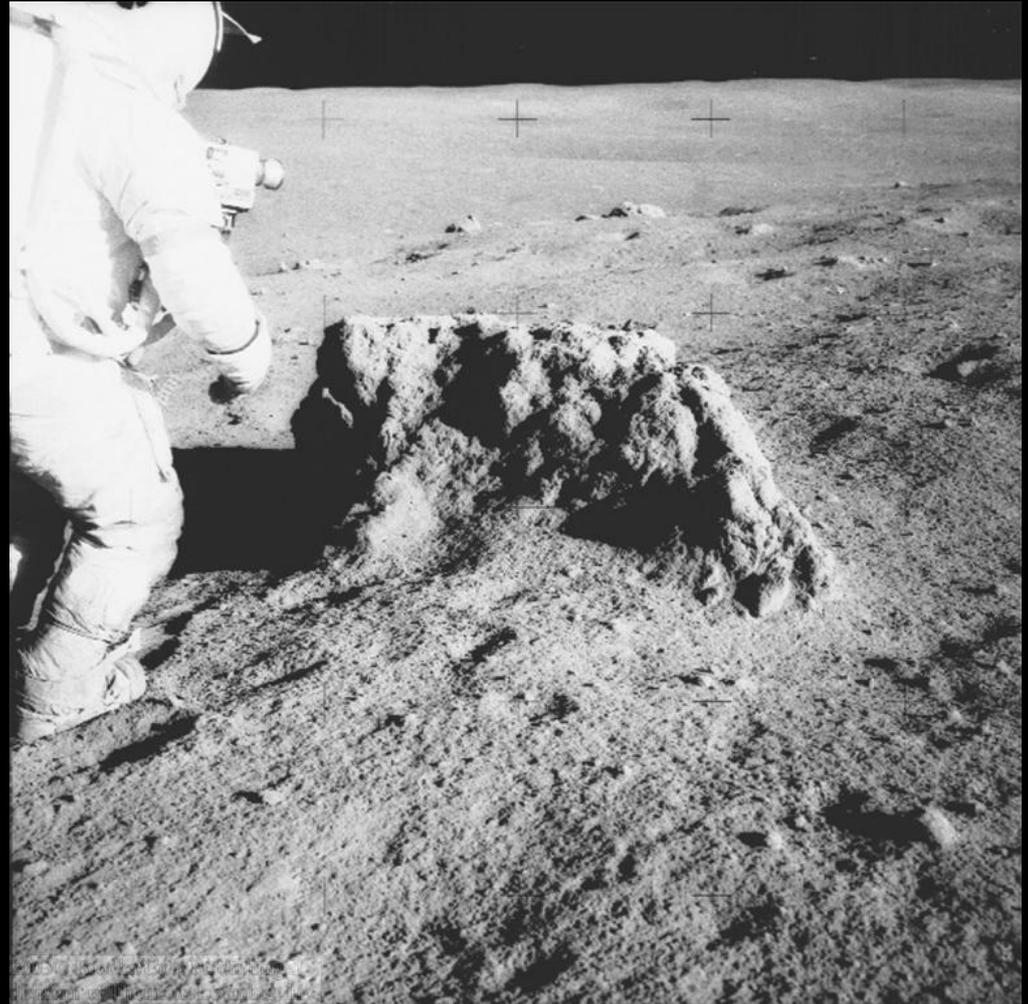
Moon must have formed in a molten state (“**sea of lava**”);

Heavy rocks sink to bottom; lighter rocks at the surface

No magnetic field → small core with little metallic iron.

Surface solidified ~ 4.6 – 4.1 billion years ago.

Heavy meteorite bombardment for the next ~ ½ billion years.



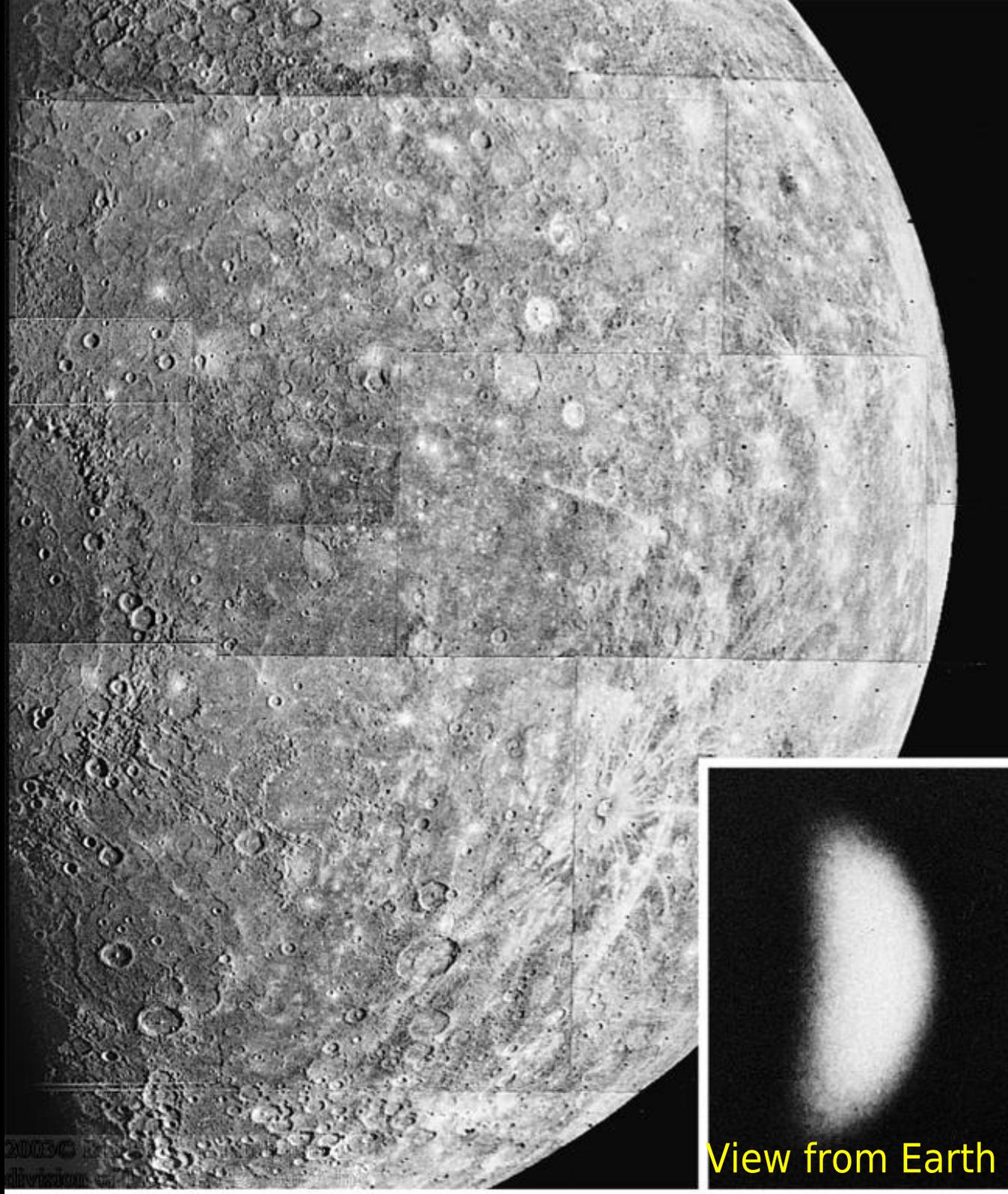
Alan Shepard (Apollo 14) analyzing a moon rock, probably ejected from a distant crater.

# Mercury

Very similar to Earth's moon in several ways:

- Small; no atmosphere
- lowlands flooded by ancient lava flows
- heavily cratered surfaces

Most of our early knowledge based on measurements by Mariner 10 spacecraft (1974 - 1975)



View from Earth

# Rotation and Revolution

Like Earth's moon, Mercury's **rotation** has been **altered by the sun's tidal forces**.

**Revolution period = 3/2 times rotation period**

Orbital period = 87.969 days

Rotation period = 58.646 days

Semi-major axis = 0.387 AU

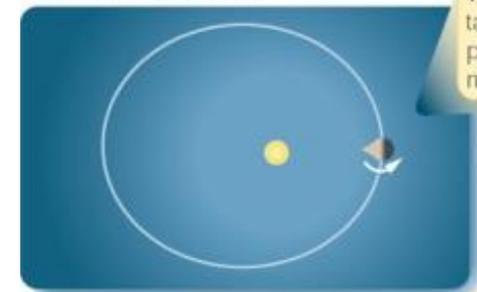
Eccentricity = 0.205

Inclination to ecliptic =  $7^\circ$

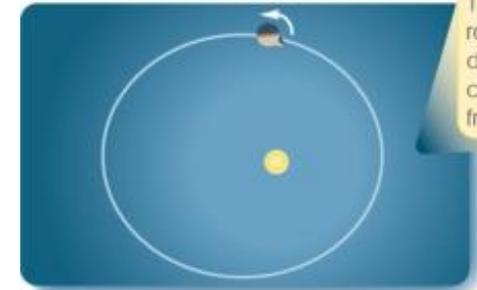
Mean radius = 2439.7 km (0.3829 Earth)

→ **Extreme day-night temperature contrast: 100 K (-173 °C) – 600 K (330 °C)**

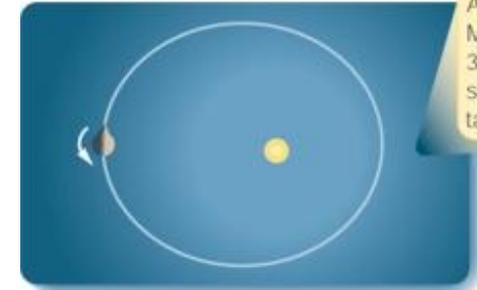
The Rotation of Mercury



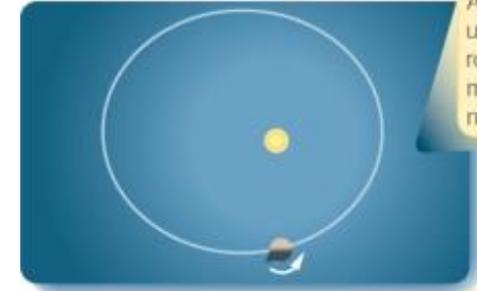
We imagine a mountain on Mercury that points at the sun. It is noon at the mountain.



The planet orbits and rotates in the same direction, counter-clockwise as seen from the north.



After half an orbit, Mercury has rotated 3/4 of a turn, and it is sunset at the mountain.



As the planet continues along its orbit, rotation carries the mountain into darkness.

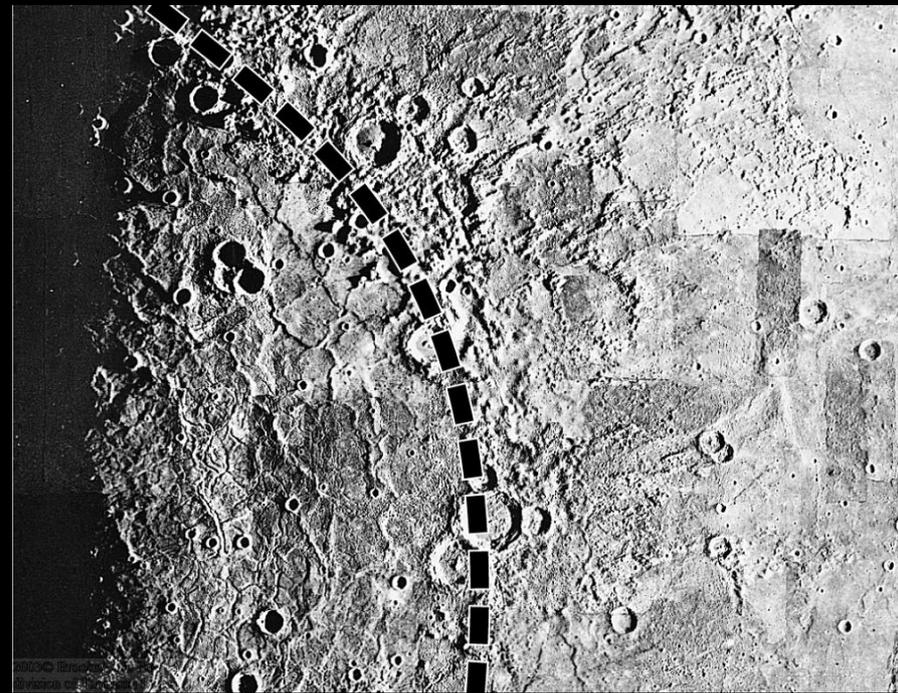
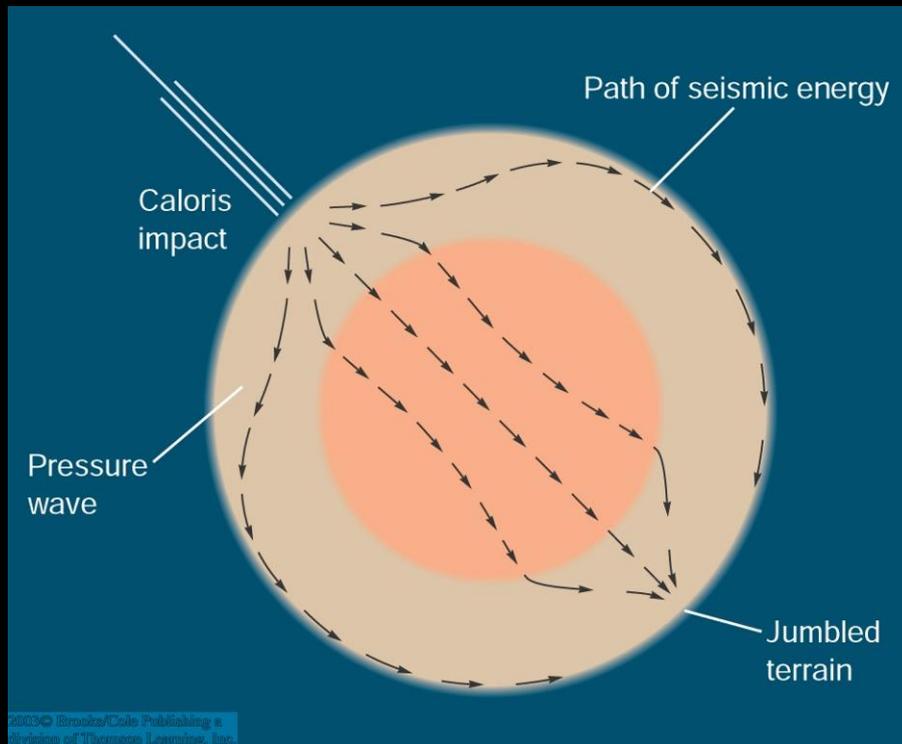


After one orbit, Mercury has rotated 1.5 times, and it is midnight at the mountain.

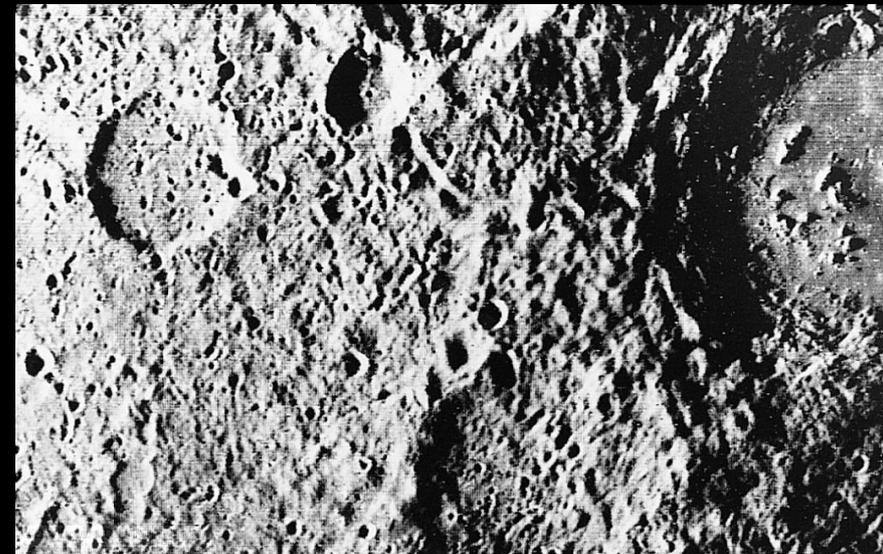
# The Surface of Mercury

Very similar to Earth's moon:  
Heavily battered with craters,  
including some large basins.

Largest basin: **Caloris Basin**



Terrain on the opposite side  
jumbled by seismic waves from  
the impact.



# Lobate Scarps



Curved cliffs, probably formed when Mercury shrunk while cooling down

# The Interior of Mercury

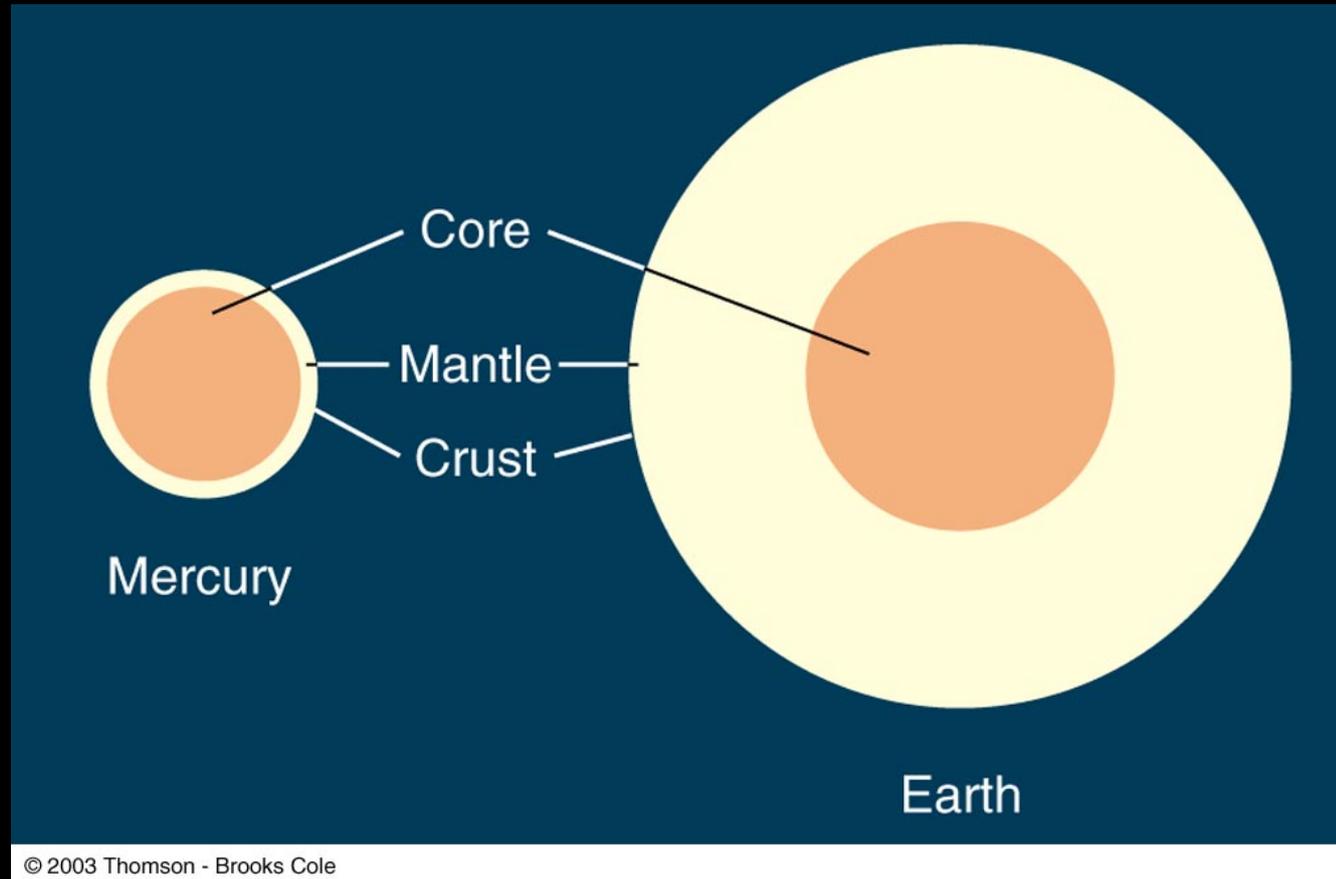
Large, metallic core.

Over 60 % denser than  
Earth's moon

Magnetic field only ~  
0.5 % of Earth's  
magnetic field.

Liquid metallic core  
should produce larger  
magnetic field.

Solid core should  
produce weaker field.



# Mercury's Atmosphere

Like the Earth's Moon, Mercury has a very volatile atmosphere. What little atmosphere exists is made up of atoms or ions blasted off its surface by the solar wind and has less than a million-billionths the pressure of Earth's atmosphere at sea level. It is composed chiefly of oxygen, sodium, and helium.

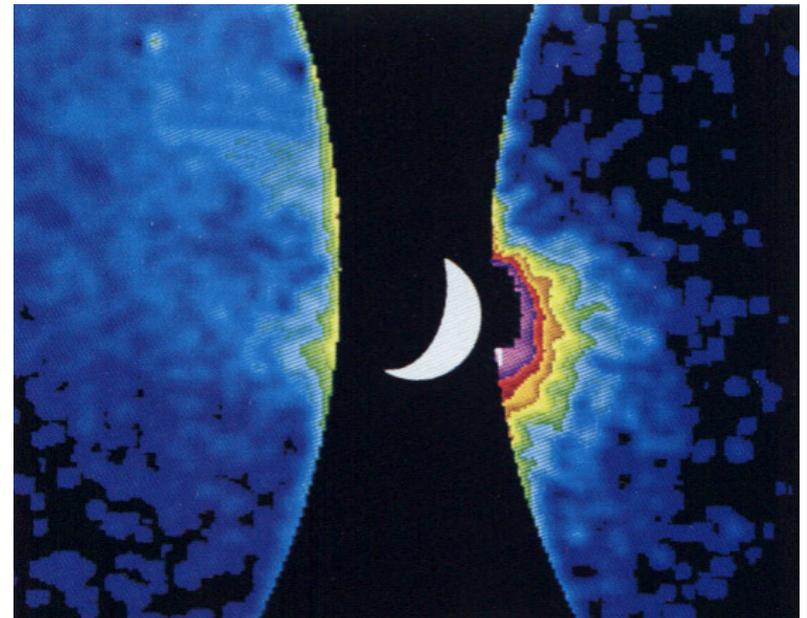
Mercury's extreme surface temperature enhances the escape of these volatile atoms into space.

With no atmosphere or hydrosphere, there has been no erosion from wind or water.

Mercury may have water ice at its north and south poles. The ice exists inside deep craters. The floors of these craters remain in perpetual shadow, so the Sun cannot melt the ice.

**Two Tenuous Atmospheres**

Species	Mercury (atoms/cm <sup>3</sup> )	Moon (atoms/cm <sup>3</sup> )
Hydrogen (H)	200	< 17
Helium (He)	6,000	2,000–4,000
Oxygen (O)	< 40,000	< 500
Sodium (Na)	20,000	70
Potassium (K)	500	16
Argon (Ar)	< 3 × 10 <sup>7</sup>	4 × 10 <sup>4</sup>





Principal Investigator: Dr. Sean C. Solomon  
Carnegie Inst. of Washington

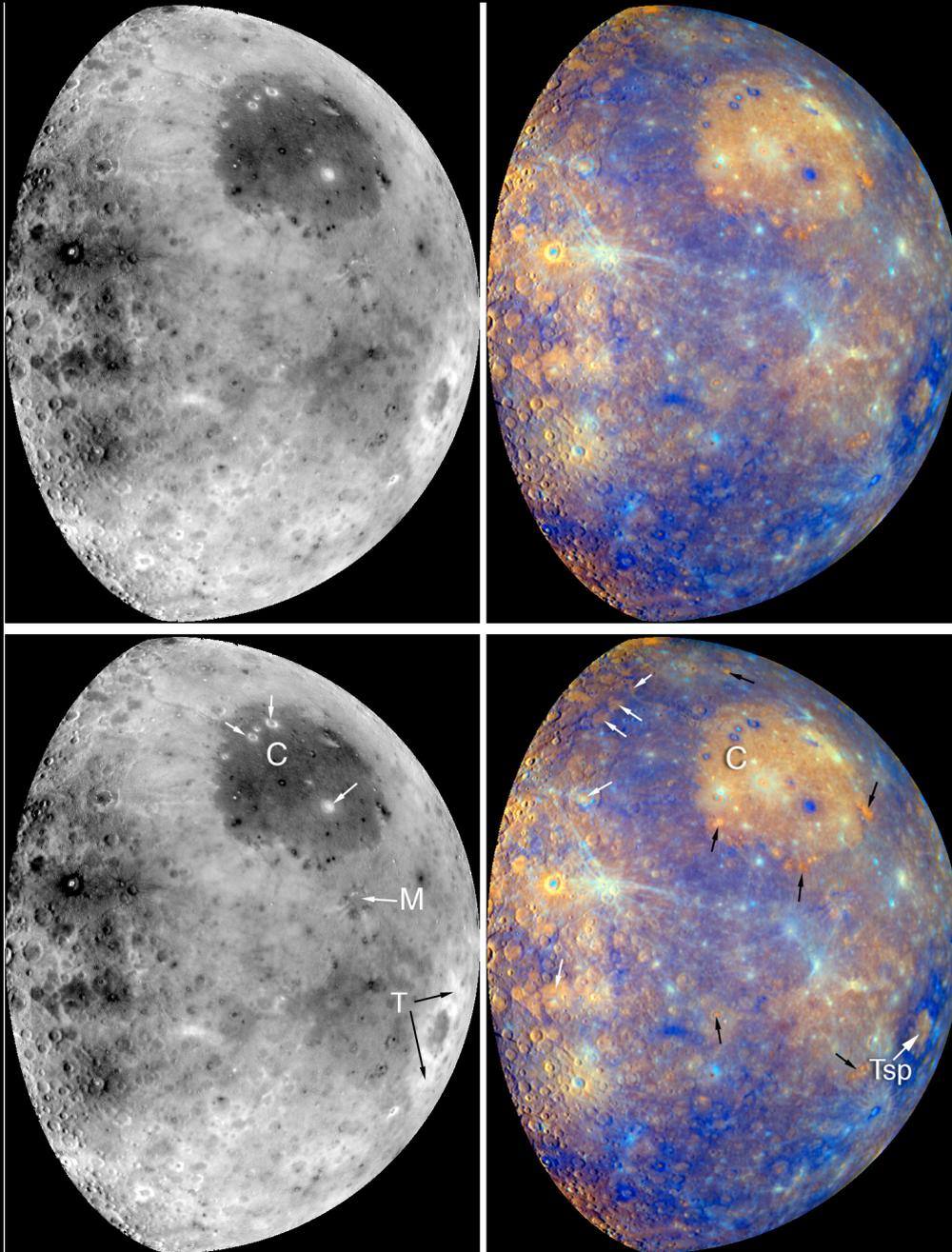
Project Management: Johns Hopkins University Applied  
Physics Laboratory

Instruments: JHU/APL, GSFC, Univ. Colorado,  
Univ. Michigan

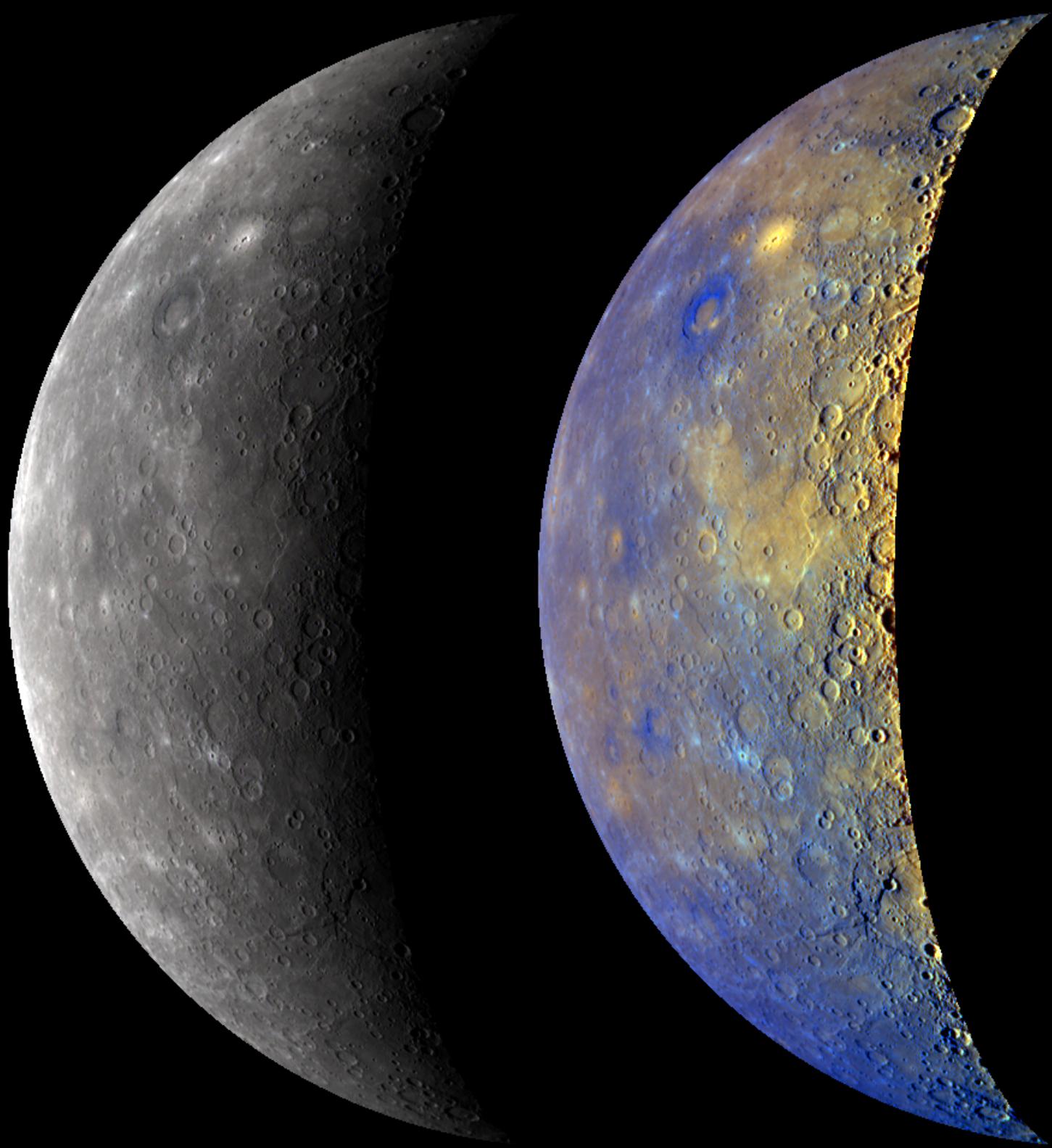
Structure: Composite Optics, Inc.

Propulsion: GenCorp Aerojet

Navigation: Jet Propulsion Laboratory



MESSENGER's WAC multi-spectral images to study compositional variations across the surface of Mercury. The white arrows identify areas of Mercury's surface that are interpreted to be relatively young volcanic plains, and the black arrows point to reddish areas interpreted to be volcanoes. Most of the color differences studied here are believed to indicate variations in the mineral composition and physical state of the rocks at different places on Mercury.



The moment of inertia factor is a dimensionless quantity that characterizes the radial distribution of mass inside a planet or satellite.

The moment of inertia of a homogenous sphere is

$$I = \frac{2}{5} MR^2 \text{ kg m}^2$$

$$\Rightarrow \frac{I}{MR^2} = 0.4$$

For Earth,  $\frac{I}{MR^2} = 0.331$ , indicative of an iron core.

For the Moon,  $\frac{I}{MR^2} = 0.3932$ , suggesting an almost uniform density.

The mean density of the Moon is  $3.4 \text{ g cm}^{-3}$ , compared to  $5.5 \text{ g cm}^{-3}$  for Earth.

Mercury: Has a tenuous atmosphere with a pressure of  $10^{-12}$  bar.

The average density of Mercury is  $5.43 \text{ g cm}^{-3}$  and the moment of inertia factor is

$$\frac{I}{MR^2} = 0.353, \text{ indicative of a liquid outer core.}$$

# **Importance of the Moon and Mercury**

- **The surface history of most other terrestrial bodies has been “lost”.**
- **The Moon is a critical laboratory in the study of early Earth history and hence the formation of a habitable planet.**
- **The Moon and Mercury surfaces provide a history of early bombardment including timing, size distribution, and composition of impacting objects.**
- **Impacts and geology of Mercury provide constraints on volatile delivery.**